The influence of tides on the sub-Amery Ice Shelf thermohaline circulation: An application of the Princeton Ocean Model to an ice shelf cavity

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Antarctica



The Amery Ice Shelf front





The ice-pump mechanism



A Jade Iceberg



Jade iceberg - Australian Antarctic Division photo by K.Sheridan © Commonwealth of Australia. May be reproduced only for non-commercial educational purposes.

Why are we interested in the Cavity beneath the Amery Ice Shelf?

- The circulation controls the rates of melting and freezing underneath the ice shelf.
- The circulation affects water mass formation in Prydz Bay and beyond
- Melting is an important way in which Antarctic glaciers lose mass (the other is by calving at the ice front).
- Removal of an ice shelf by melting and/or calving may affect the rates at which Antarctic glaciers flow, and hence affect the mass balance of ice over the continent.
- How are the above affected by global warming, both in the past, and in the future?

Models

- Determann/Gerdes (AWI), 1994: Idealised cavities
- Grosfeld/Gerdes (AWI/Bremen), 1997: Idealised cavities and Filchner-Ronne
- Gerdes/Determann/Grosfeld (AWI), 1999: Filchner-Ronne
- Grosfeld/Sandhager/Lange (Bremen/AWI/Munster), 2001: Idealised cavity (→ Larsen + Flichner-Ronne) coupled with ice model
- Beckmann/Hellmer (AWI): Whole Southern Ocean (BRIOS2.2)
- Holland/Jenkins (New York/BAS):
 Filchner-Ronne and Ross based on MICOM model.
- Williams/Warner/Budd (Antarctic CRC), 1998-2002: Amery (Gerdes/Determann/Grosfeld model) - without tides.
- Hunter/Hemer (Antarctic CRC): Amery (based on POM) - with and without tides.

The Ice-Water Interface



3-Equation Method:

(~Holland and Jenkins, 1999)

Freezing point:

$$T_B = T_B(S_B, P)$$

Heat:

$$Q_T^W = (E_T(T_M - T_B) + UT_B)\rho C_p$$
$$= Q_T^I - U\rho L$$

Salt:

$$Q_S^W = (E_S(S_M - S_B) + US_B)\rho$$
$$= 0$$

Bed elevation (\bullet) and Ice Draft (all) measurement locations.



Zonal Mean Bed Elevation and Ice Draft of Amery Ice Shelf Cavity. (Dashed - A; Solid - B; Dash-Dot - C; Dotted - CADA)



Bathymetry



Ice Shelf Draft



Modifications to POM for the Ice Shelf application:

- Define additional horizontal U, V and E masks for ice shelf
- Apply surface pressure to depress ocean surface by D (i.e. apply adjustment D to all heights that are multiplied by g, except in baroclinic pressure and buoyancy gradients).
- Dynamics at ice/ocean interface: invert bottom friction code (yields u*).
- Thermodynamics at ice/ocean interface: (2-equation formulation of Holland and Jenkins, 1999 removes third (salt) equation: $S_M = S_B$)
 - Interface assumed to be at local freezing temperature (T_f) (converted to potential temperature).
 - Heat flux (wtsurf) assumed $\propto u^*(T_{ocean} T_f)$.
 - Heat flux yields melting or freezing rate. $vflux \propto wtsurf$
 - Melting (freezing) leads to addition (removal) of freshwater at in-situ freezing temperature (i.e. Additional heat and salt flux term).
 - Add source/sink of water to volume conservation equation in external mode.
 - Set w(i, j, kb) = 0, reverse order of vertical integration of w and stop at k = 2 (puts source/sink of water into top cell).

Initial and Open BC's

- Initial and Open Boundary conditions based on Winter conditions (deep mixing) in Prydz Bay.
 - Salinity of 34.5, Temperature at surface freezing temperature $(-1.8976^{\circ}C)$ - uniform throughout the water column.
 - S and T are "cobbled" at open boundaries.
- Tidal elevations at OB's are determined from tidal constituents $(M_2, S_2, K_1, and O_1)$ interpolated onto OB gridcells from CADA tide model (Padman et al., 2002).

Runs include:

- Tides switched on and off 4km and 8km resolution.
- Other initial temperatures (simulating global warming scenarios):
 - $T_f(surface) + 0.2^{\circ}C$
 - $T_f(surface) + 1^\circ C$



Hence, spin up time is tens of years

Longitudinal Density Section (σ_{θ}) :







Zonally-Integrated Streamfunction (Sv): No Tides (Circulation clockwise around +ve features)

Vertically-Integrated Streamfunction (Sv): No Tides (Circulation clockwise around +ve features)





Freezing rate (ma^{-1}) : No Tides (Positive is freezing; bold separates freezing/melting)



'Observed' Freezing/Melting



Blue: 10 m a^{-1} freezing Red: 10 m a^{-1} melting

Kinematic estimation based on SAR, altimetry and ice density model

(Young, Hyland and Gale)

Vertically-Integrated Streamfunction (Sv): No Tides; $+1^{\circ}C$ (Circulation clockwise around +ve features)



Freezing rate (ma^{-1}) : No Tides; $+1^{\circ}C$ (Positive is freezing; bold separates freezing/melting)





Average slope = $43Gta^{-1\circ}C^{-1}$ (c.f. $25Gta^{-1\circ}C^{-1}$; Williams et al., 2002) An increase in mass loss of $40Gta^{-1}$ could remove the ice shelf in 1000 years

Future Work

- Coupled cavity/ice shelf model (JH, Hobart)
- Upgraded ice draft from hydrostatics/altimetry and radar (JH, Hobart)
- Incorporation of frazil ice (JH, Hobart)
- Investigation of annual cycle (requires a sea-ice model or prescribed surface fluxes in open ocean). (JH, Hobart; MH Galway)
- Model Intercomparison study for idealised ice-shelf domains (MH Galway)
 - POM; Hemer NUI, Galway, Ireland, Hunter ACE CRC, Hobart
 - Generalised coordinate Cox-Bryan model; Williams NIWA, NZ
 - MICOM; Holland, New York
 - SPEM/SCRUM??; AWI

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Personnel:

- 3 Permanent Staff
- 2 Post-Docs (as of Sept)
- 6 Post-grad students

Projects:

- Hydrodynamic model focussed on the North-East Atlantic - Irish territorial waters. (POM -> ROMS??)
- Operational forecasting model of the Irish Sea (POM or POLCOMS)
- Flushing Studies of Irish Coastal Waters (POM)
- Kinetics and transport of scallop and sea-jelly plankton larvae (POM)